

WG2 Doppler lidar



Report

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Updates since Varna

- Instrument characterization
 - Methods for improving data quality and sensitivity

Instrument retrievals

- Wind
- Wind gusts
- Turbulence
- BL classification



Updates since Varna

- Obtaining turbulent retrievals from Leosphere instrument
 - Standard retrieval from vertical stare
 - VAD retrieval from scans at multiple elevations
 - Shu Yang presentation tomorrow



Wind gusts – paper published

Suomi, I., Gryning, S.-E., O'Connor, E. J. and Vihma, T. (2017), Methodology for obtaining wind gusts using Doppler lidar. Q.J.R. Meteorol. Soc., 143: 2061–2072. doi:10.1002/qj.3059





LLJs – paper in press

Suomi, I., Gryning, S.-E., O'Connor, E. J. and Vihma, T. (2017), Methodology for obtaining wind gusts using Doppler lidar. Q.J.R. Meteorol. Soc., 143: 2061–2072. doi:10.1002/qj.3059





Boundary layer classification – more updates

Fields: Turbulent source, and in-contact-with Includes: LLJ





Boundary layer classification – more updates

Fields: Turbulent source, and in-contact-with Includes: LLJ





SWG and 3 STSMs

- SWG for winds, turbulence, ...
- STSM Shu Yang, to FMI
 - Leosphere characterization
- STSM Umar Saeed, to FMI
 - Combining DWL + MWR
- STSM Ekaterina Batchvarova, to DTU
 - DWL forecast verification



SWG on winds, turbulence

Recommendations for winds

- Retrieval methods
- Scan selection
- Paper plans
 - SOPs for siting, operation, pre-processing, network
 - Wind retrieval VAD vs DBS
 - Contribution to BAMS 'TOPROF' paper



Leosphere background correction - STSM

Instrument usually stable, but occasional jumps





Leosphere background correction - STSM

Apply 'similar' correction to that applied to Halo systems.

Some differences because of slightly different internal processing implementation

-	Original signal
1	Cloud screening
1.1	2D variance-based cloud screening
1.2	Final cloud screening utilising Cook's distance
1.3	Calculation of the background shape and refilling
V	
2	Step detection
2.1	Multilevel 1D stationary wavelet decomposition
2.2	Peak detection from the wavelet decomposition output
V	
3	Correction of step changes and background shape
3.1	Calculation of the background shape of each step
3.2	Correction for the shape and amplitude of the background
V	
4	Removal of possible remnant outlier profiles (optional)
4.1	Detection of outlier profiles from corrected background
4.2	Removal of the detected profiles
•	Background correction of signal
•	Recalculation of attenuated backscatter coefficients



Leosphere background correction





Leosphere background correction





'High-resolution' MWR retrieval

• Identify stable, neutral, unstable



Figure1. Comparison of MWR and radiosonde temperature profiles under convective-driven mixing



'High-resolution' MWR retrieval

• Identify stable, neutral, unstable



















SCIENTIFIC REPORT



- ACTION: ES1303 TOPROF
- STSM: COST-STSM-ES1303-38266
- TOPIC: The use of lidar measurements for evaluation of wind-speed prediction by numerical models
- VENUE: DTU Wind Energy, Denmark
- PERIOD: 08 August 16 August, 2017
- •
- Host: Jake Badger (DTU Wind Energy, Denmark)
- Applicant: Ekaterina Batchvarova (NIMH-BAS, Bulgaria)
- Submission date: 03.09.2017

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• Contribution by: Ekaterina Batchvarova (NIMH-BAS, Bulgaria) and Sven-Erik Gryning (DTU, Denmark)







Data availability for different CNR values



For the same CNR more data are available in the marine air compared to the air over land

The WRF simulation – courtesy to Roger Floors, DTU

The model data set is created with the Weather Research and Forecast model WRF (Skamarock et al, 2004) in analysis mode; FNL global boundary conditions available every 6 hours on a 1° x 1° grid; two nested domains of horizontal grid size of 18 and 2 km; Noah land surface scheme (Chen and Dudhia 2001), MYNN surface layer scheme (Nakanishi and Niino 2009), Thompson microphysics scheme (Thompson et al. 2004), and the 1.5 order closure Mellor-Yamada Nakanishi and Niino level 2.5 (MYNN, Nakanishi and Niino (2009) planetary boundary-layer (PBL) scheme.

The WRF model was configured to calculate the meteorological parameters at 41 vertical levels from the surface to pressure level 100 hPa. Eight of these levels were within the height range of 600 m and the first model level was at ~14 m. The simulations were initialized every 10 days at 12:00 GMT and after a spin up of 24 hours a time series of 10-min output was picked out from the simulated meteorological data from hour 25 to 264. In order to prevent the model from

Applying CNR=-22 dB leads to higher mean wind speeds compared to CNR=-35 dB WRF is always underestimating the wind speed at FINO3 For CNR=-22 dB the difference is smaller compared to CNR=-35 dB



Different k-parameters profiles over land and over sea Smaller variability /bigger k/ for CNR=-35 dB compared to CNR=-22 dB WRF overestimates k or underestimates variability



The outcome of the study can be summarized as:

In general, WRF underestimates the wind speed and overestimates the Weibull shape parameter at all levels, which means that the model suggests lower values and lower variability for the wind speed at all levels up to 600 m.

Thus, when comparing all WRF data to lidar data with strong CNR filter applied, the underestimation will be bigger than presented here.

Also, if high quality lidar data are assimilated into WRF, there will be shift towards higher wind speeds, which may reduce the difference between model and observations.